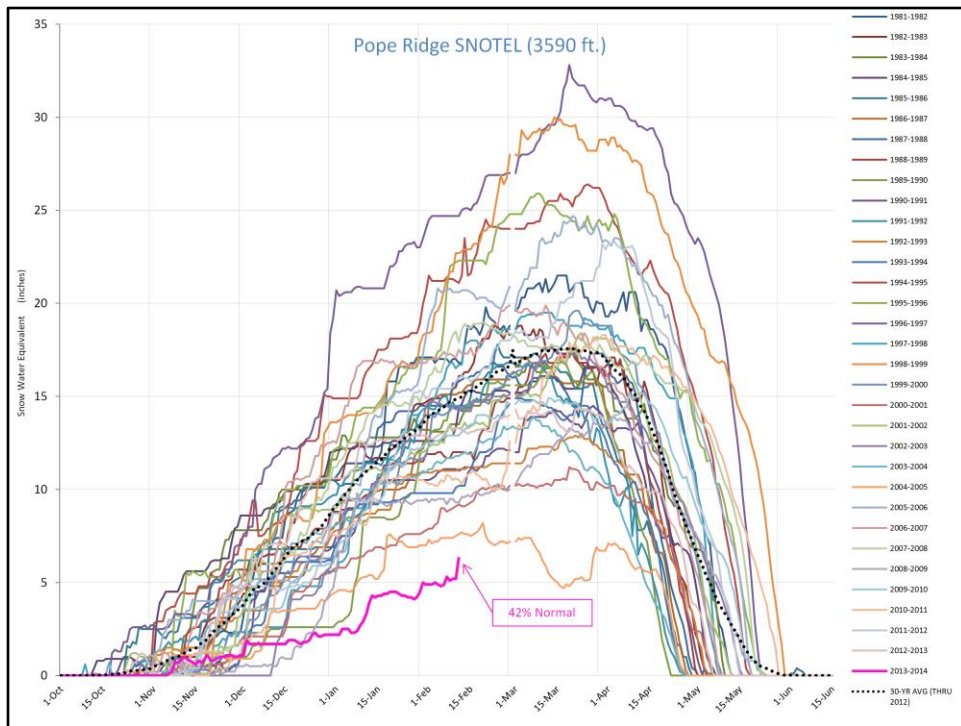
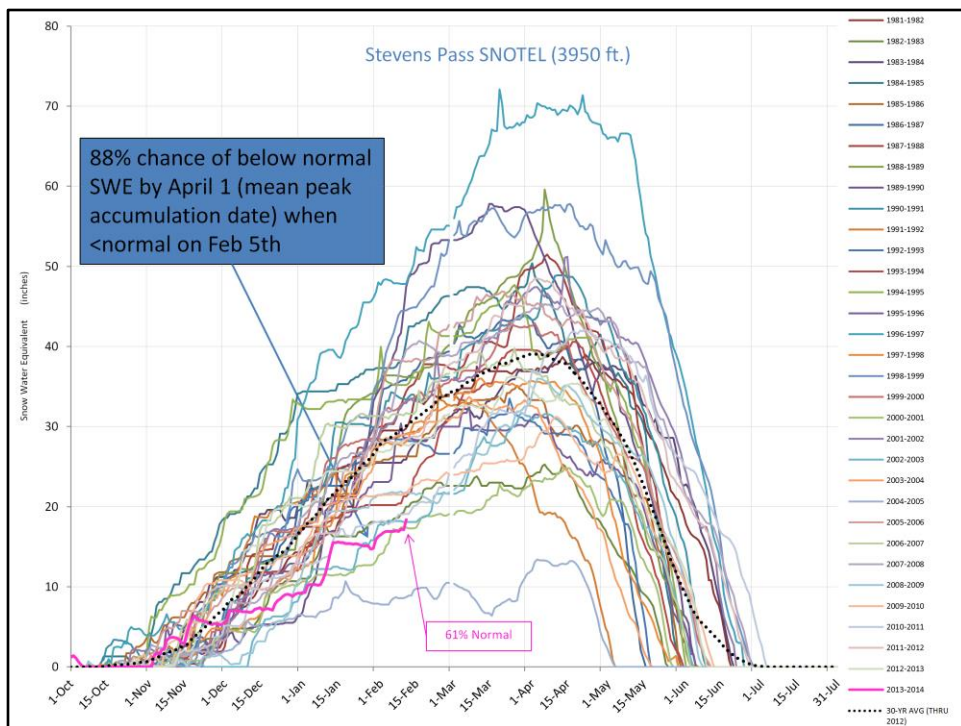
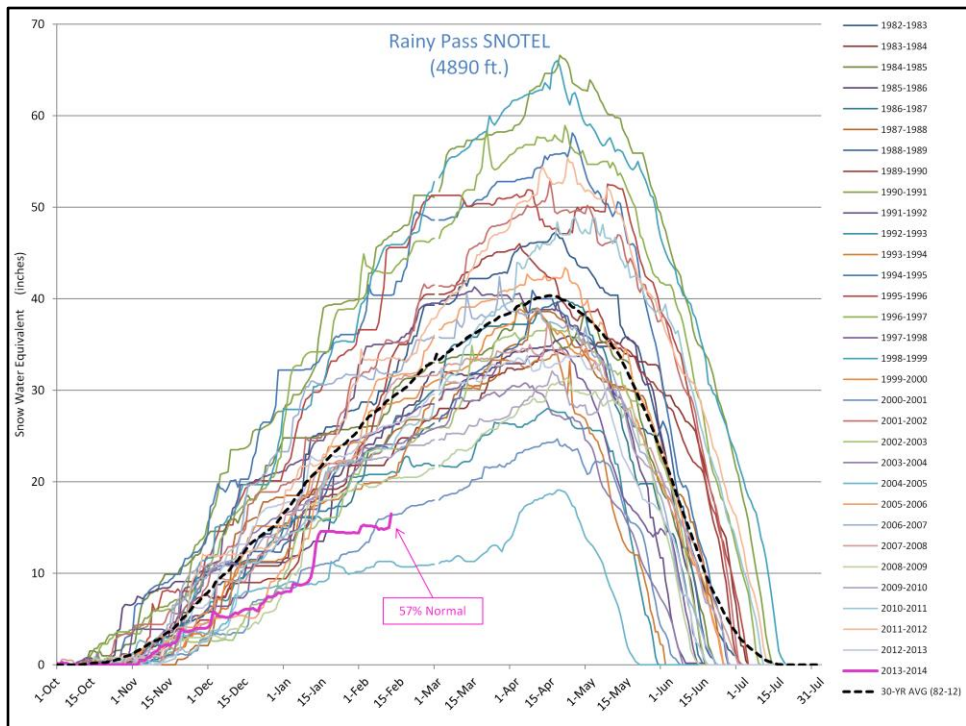


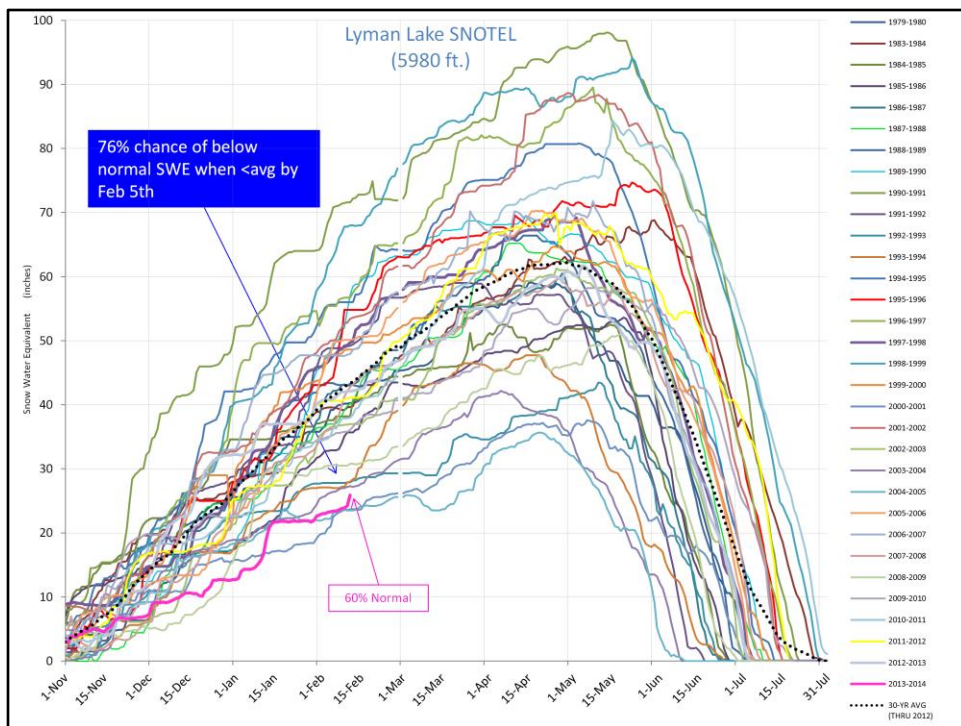
Upper Columbia SNOTEL Data: Current trends and our seasonal outlook

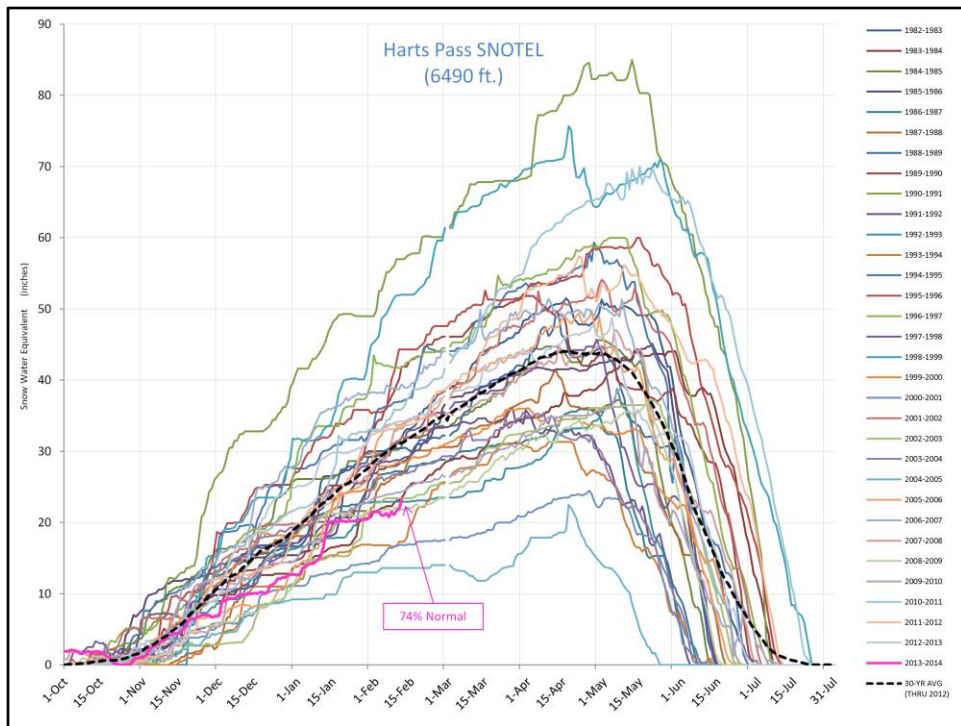














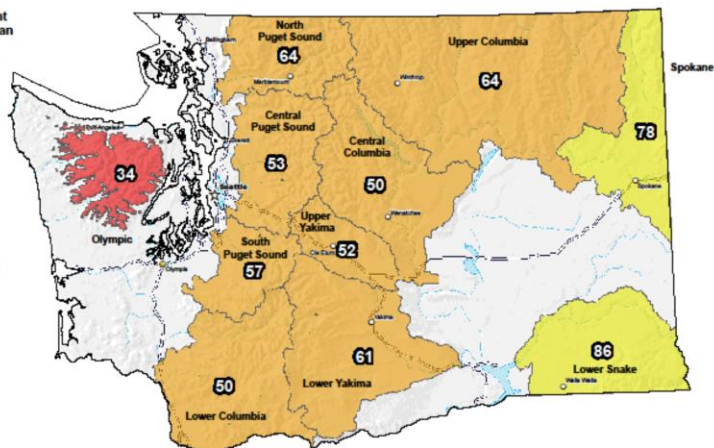






Feb 05, 2014

	unavailable *
	<50%
	50 - 69%
	70 - 89%
	90 - 109%
	110 - 129%
	130 - 149%
	>=150%



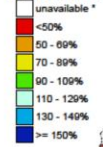
The snow water equivalent percent of normal represents the current snow water equivalent found at selected SNOTEL sites in or near the basin compared to the average value for those sites on this day. Data based on the first reading of the day (typically 00:00).

Prepared by the USDA/NRCS National Water and Climate Center
Portland, Oregon <http://www.wcc.nrcs.usda.gov/igis/>
Based on data from <http://www.wcc.nrcs.usda.gov/reports/>
Science contact: Jim.Marron@por.usda.gov 503 414 3047

Westwide SNOTEL Current Snow Water Equivalent (SWE) % of Normal

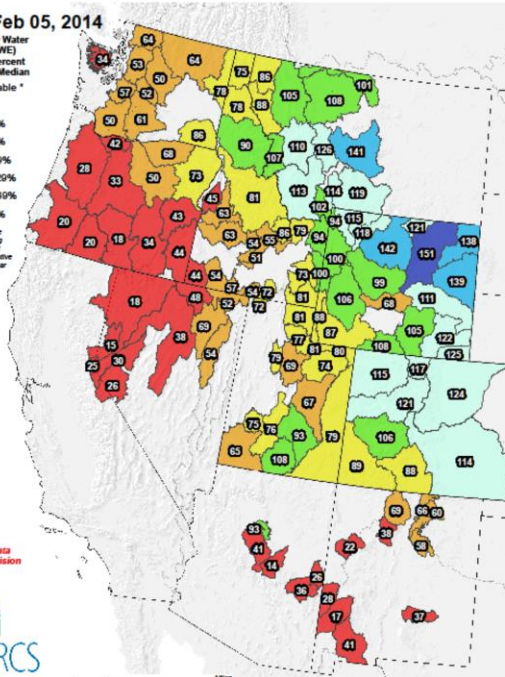
Feb 05, 2014

Current Snow Water
Equivalent (SWE)
Basin-wide Percent
of 1981-2010 Median



* Data unavailable
at time of posting
or measurement
is not representative
at this time of year

Provisional data
subject to revision



Snow Water Equivalents (inches)

Provided by the California Cooperative Snow Surveys

Data For: 11-Feb-2014

% Apr 1 Avg. / % Normal for this Date



Change Date :



11-Feb-2014

Refresh Data

NORTH

Data For: 11-Feb-2014

Number of Stations Reporting 28
Average snow water equivalent 3.9"
Percent of April 1 Average 14%
Percent of normal for this date 19%

CENTRAL

Data For: 11-Feb-2014

Number of Stations Reporting 44
Average snow water equivalent 7.9"
Percent of April 1 Average 26%
Percent of normal for this date 36%

SOUTH

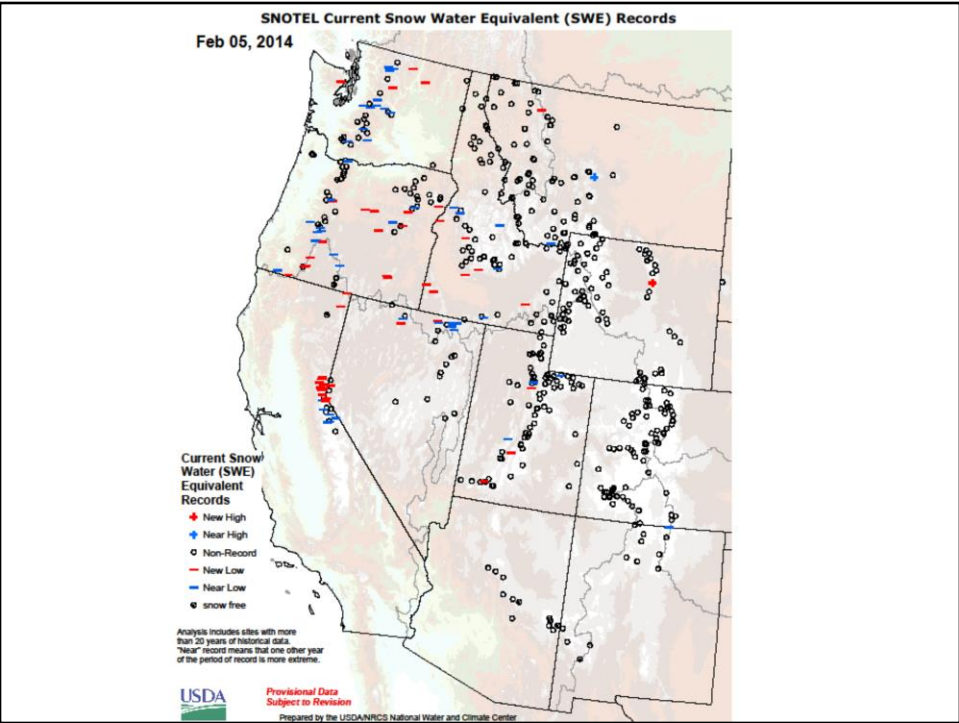
Data For: 11-Feb-2014

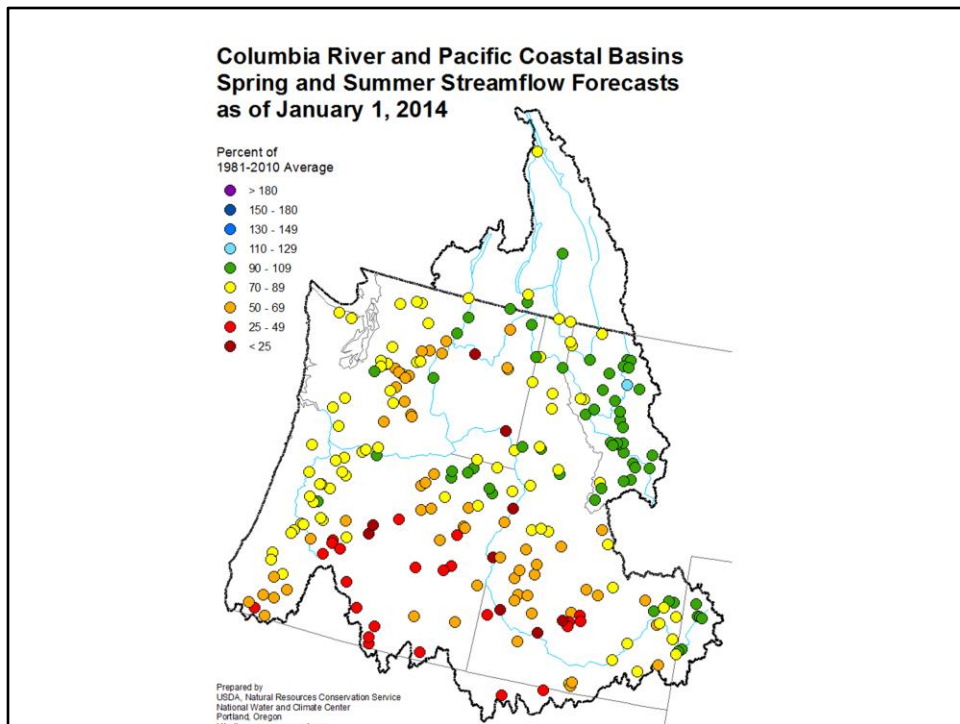
Number of Stations Reporting 33
Average snow water equivalent 4.7"
Percent of April 1 Average 18%
Percent of normal for this date 26%

STATEWIDE SUMMARY

Data For: 11-Feb-2014

Number of Stations Reporting 105
Average snow water equivalent 5.8"
Percent of April 1 Average 20%
Percent of normal for this date 28%

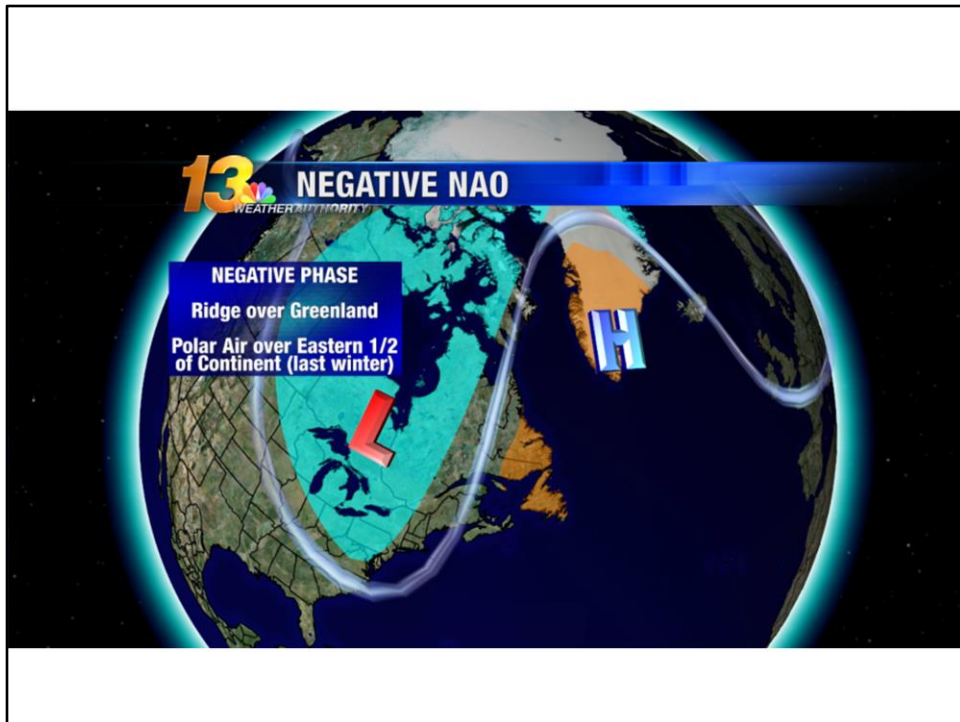




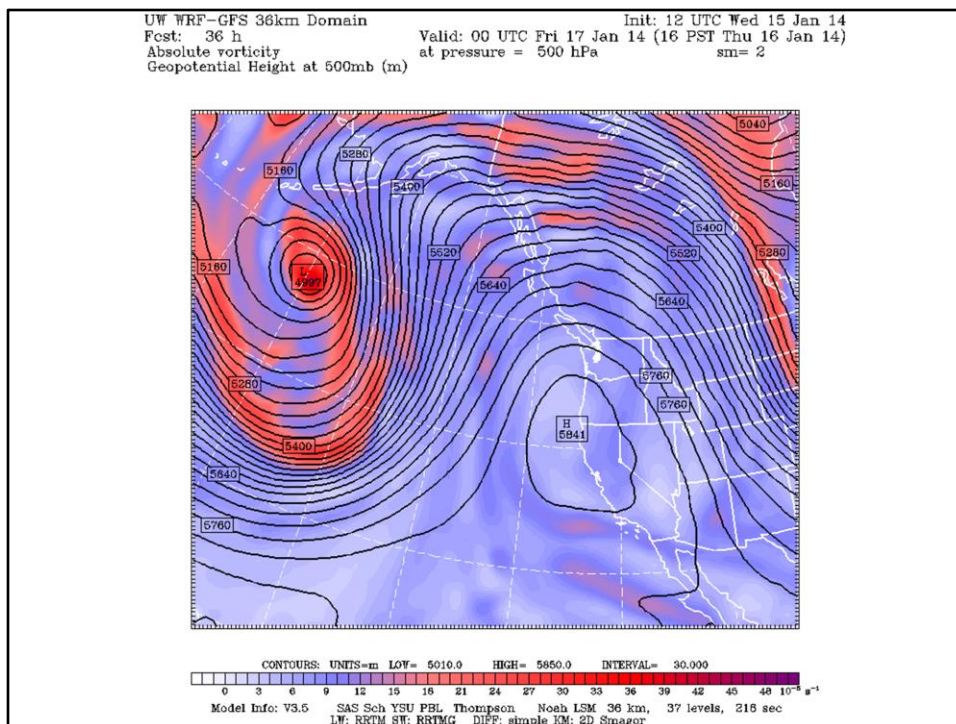
Plot showing the 50 Percent Chance of Exceedance Forecast. There is a 50 percent chance that the actual streamflow volume will exceed this forecast value, and there is a 50 percent chance that the actual streamflow volume will be less than this forecast value. Generally, this forecast is the middle of the range of possible streamflow volumes that can be produced given current conditions.



This shows the persistent pattern we have observed for most of this winter so far.



The ridge in the western US has led to strongly elongated low pressure in the eastern US with another downstream blocking high pressure over Greenland.

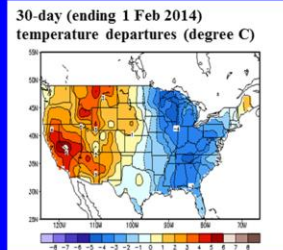
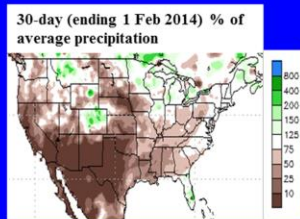


Typical eastern Pacific surface plot for an Omega Block pattern.

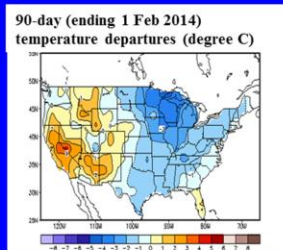
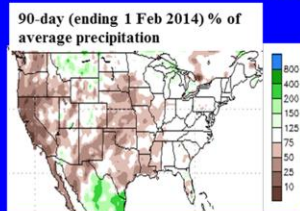


U.S. Temperature and Precipitation Departures During the Last 30 and 90 Days

Last 30 Days

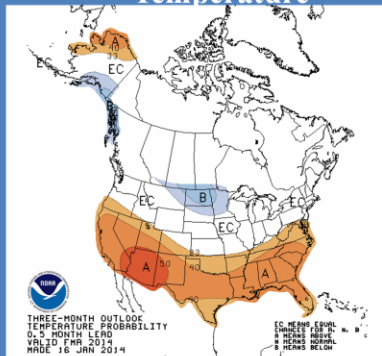


Last 90 Days

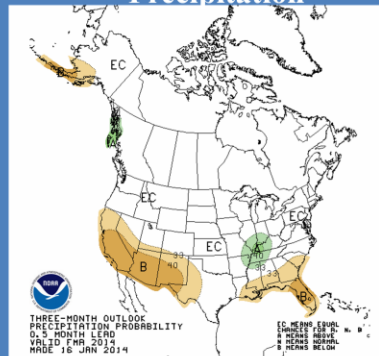


U. S. Seasonal Outlooks February – April 2014

Temperature



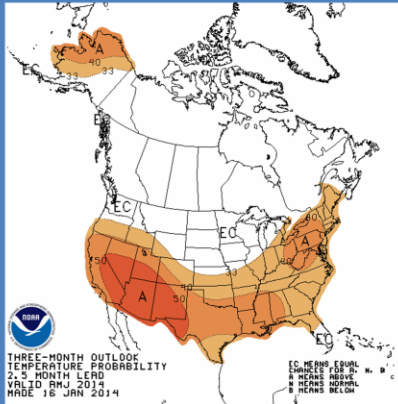
Precipitation



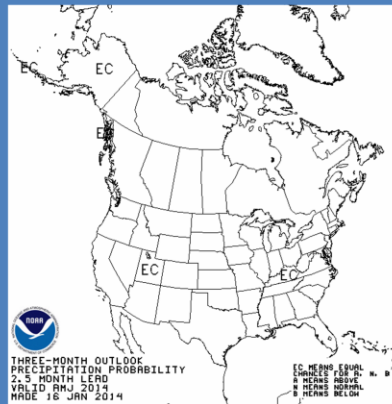
The seasonal outlooks combine the effects of long-term trends, soil moisture, and, when appropriate, ENSO.

U. S. Seasonal Outlooks April – June 2014

Temperature



Precipitation



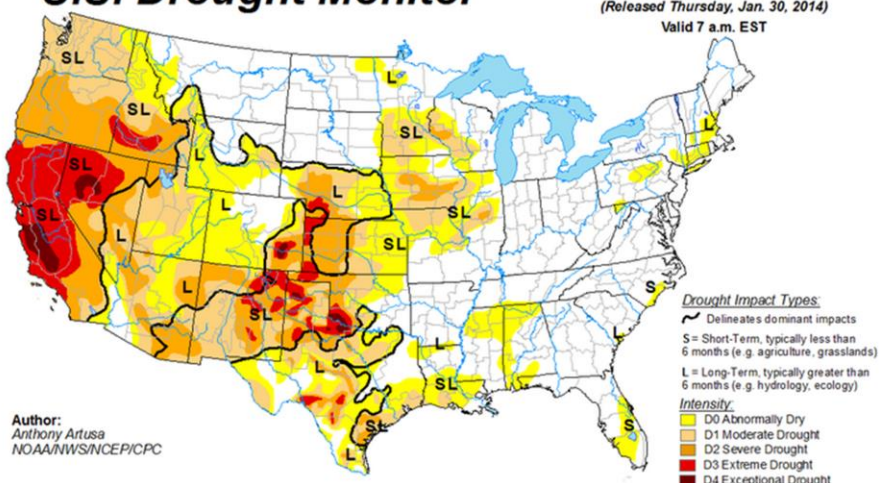
The seasonal outlooks combine the effects of long-term trends, soil moisture, and, when appropriate, ENSO.

U.S. Drought Monitor

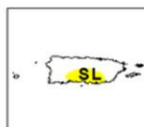
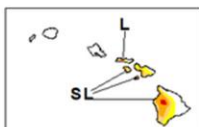
January 28, 2014

(Released Thursday, Jan. 30, 2014)

Valid 7 a.m. EST



Author:
 Anthony Artusa
 NOAA/NWS/NCEP/CPC



The Drought Monitor focuses on broad-scale conditions. Local conditions may vary. See accompanying text summary for forecast statements.



<http://droughtmonitor.unl.edu/>



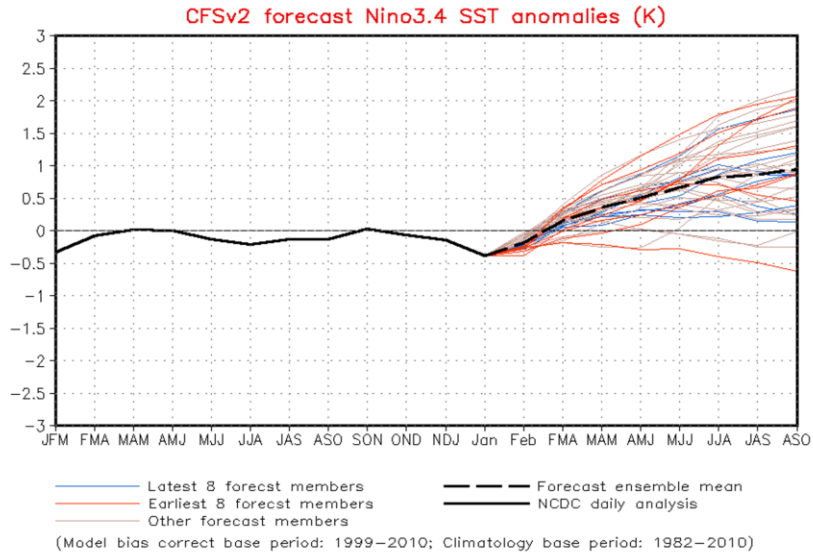
Recent Pacific warm (red) and cold (blue) episodes based on a threshold of $\pm 0.5^{\circ}\text{C}$ for the Oceanic Nino Index (ONI) [3 month running mean of ERSST.v3b SST anomalies in the Nino 3.4 region (5N-5S, 120-170W)]. For historical purposes El Niño and La Niña episodes are defined when the threshold is met for a minimum of 5 consecutive over-lapping seasons. The complete table going back to DJF 1950 can be found by clicking: [Historical ONI Values](#)

Year	DJF	JFM	FMA	MAM	AMJ	MIJ	JJA	JAS	ASO	SON	OND	NDJ
2002	-0.2	0.0	0.1	0.3	0.5	0.7	0.8	0.8	0.9	1.2	1.3	1.3
2003	1.1	0.8	0.4	0.0	-0.2	-0.1	0.2	0.4	0.4	0.4	0.4	0.3
2004	0.3	0.2	0.1	0.1	0.2	0.3	0.5	0.7	0.8	0.7	0.7	0.7
2005	0.6	0.4	0.3	0.3	0.3	0.3	0.2	0.1	0.0	-0.2	-0.5	-0.8
2006	-0.9	-0.7	-0.5	-0.3	0.0	0.1	0.2	0.3	0.5	0.8	1.0	1.0
2007	0.7	0.3	-0.1	-0.2	-0.3	-0.3	-0.4	-0.6	-0.8	-1.1	-1.2	-1.4
2008	-1.5	-1.5	-1.2	-0.9	-0.7	-0.5	-0.3	-0.2	-0.1	-0.2	-0.5	-0.7
2009	-0.8	-0.7	-0.5	-0.2	0.2	0.4	0.5	0.6	0.8	1.1	1.4	1.6
2010	1.6	1.3	1.0	0.6	0.1	-0.4	-0.9	-1.2	-1.4	-1.5	-1.5	-1.5
2011	-1.4	-1.2	-0.9	-0.6	-0.3	-0.2	-0.2	-0.4	-0.6	-0.8	-1.0	-1.0
2012	-0.9	-0.6	-0.5	-0.3	-0.2	0.0	0.1	0.4	0.5	0.6	0.2	-0.3
2013	-0.6	-0.6	-0.4	-0.2	-0.2	-0.3	-0.3	-0.3	-0.3	-0.2	-0.3	
2014												

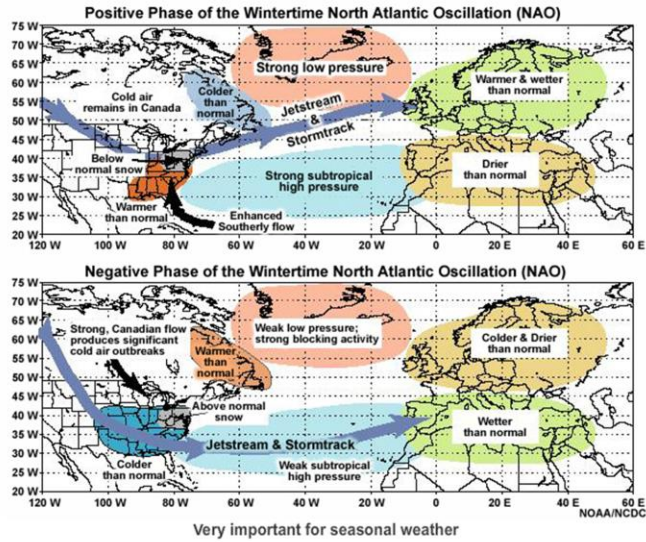
Natural, seasonal variability—
ENSO (El Niño/La Niña)

We have been in ENSO neutral conditions since early 2012.

- Most models predict ENSO-neutral (-0.5°C to $+0.5^{\circ}\text{C}$) to continue through the Northern Hemisphere spring. After that, models predict either ENSO-neutral or El Niño (greater or equal to $+0.5^{\circ}\text{C}$) during the Northern Hemisphere summer 2014.



Most models and the ensemble (black dashed line) show an El Niño developing by the April-May-June time period.



Natural, seasonal variability— North Atlantic Oscillation

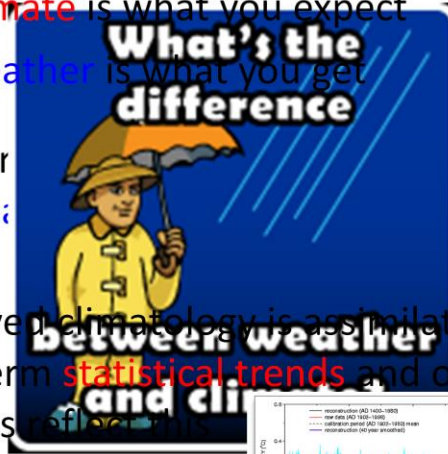
The North Atlantic Oscillation is another well-known natural, shorter duration phenomena that can affect seasonal weather patterns.

It's worth remembering the old adage.....

Climate is what you expect

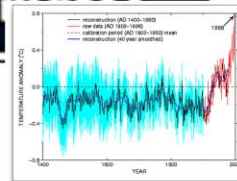
Weather is what you get

In other
and we:

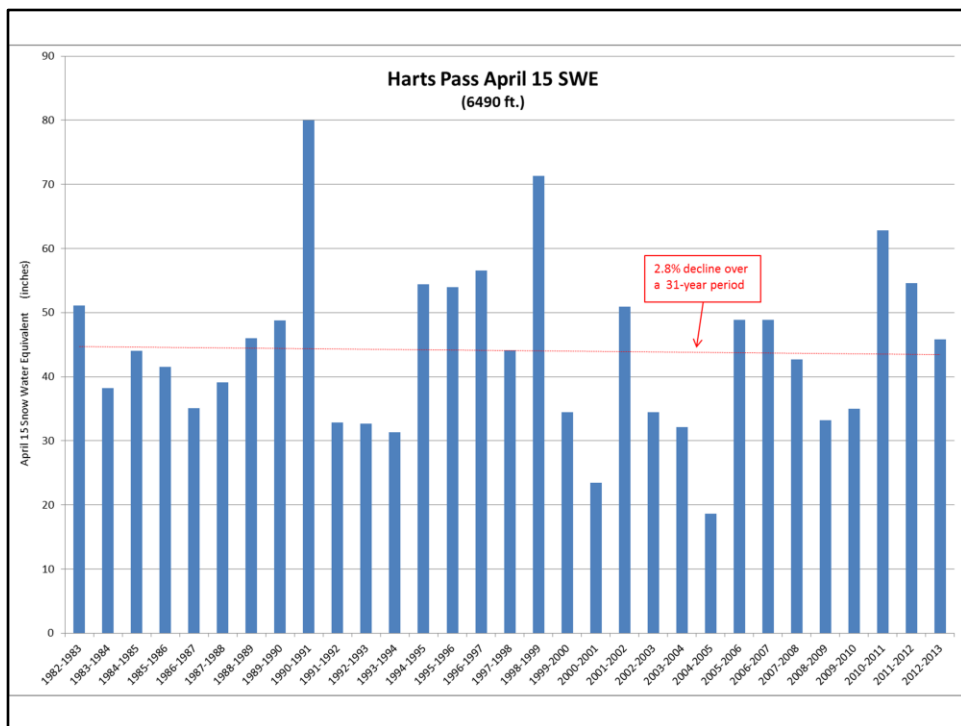


m trends
mena

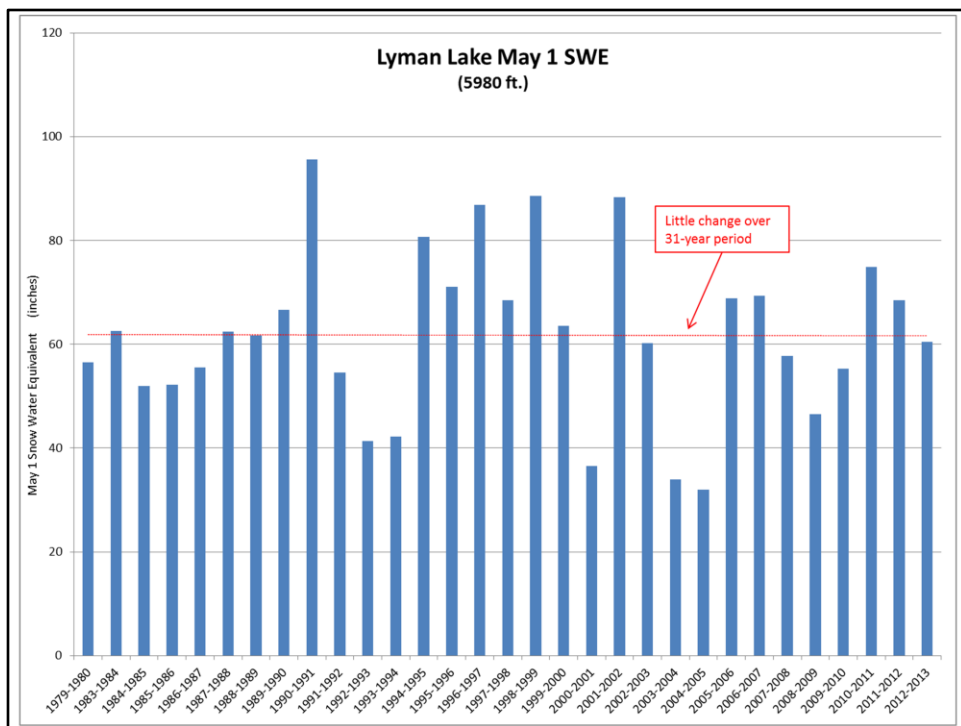
Observed climatology is assimilative of
long-term **statistical trends** and our SNOTEL
stations reflect



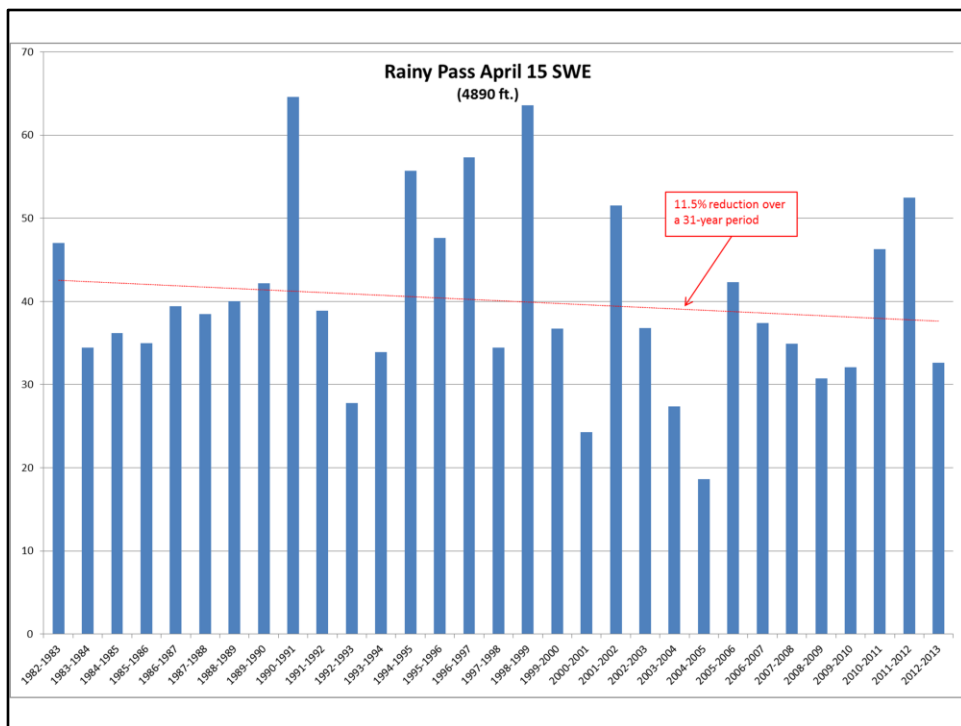
Folks need to understand the difference between weather and climate—they are different but commonly misunderstood.



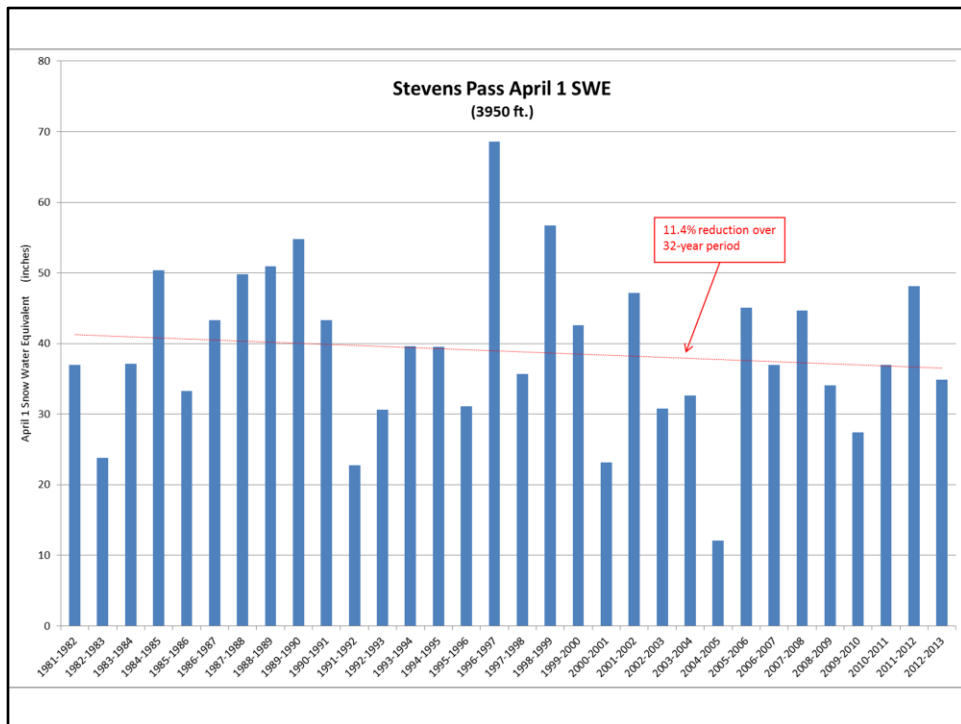
One of our high-elevation sites that shows little decline in April 15 Snow Water Equivalent over the 30-year timeframe.



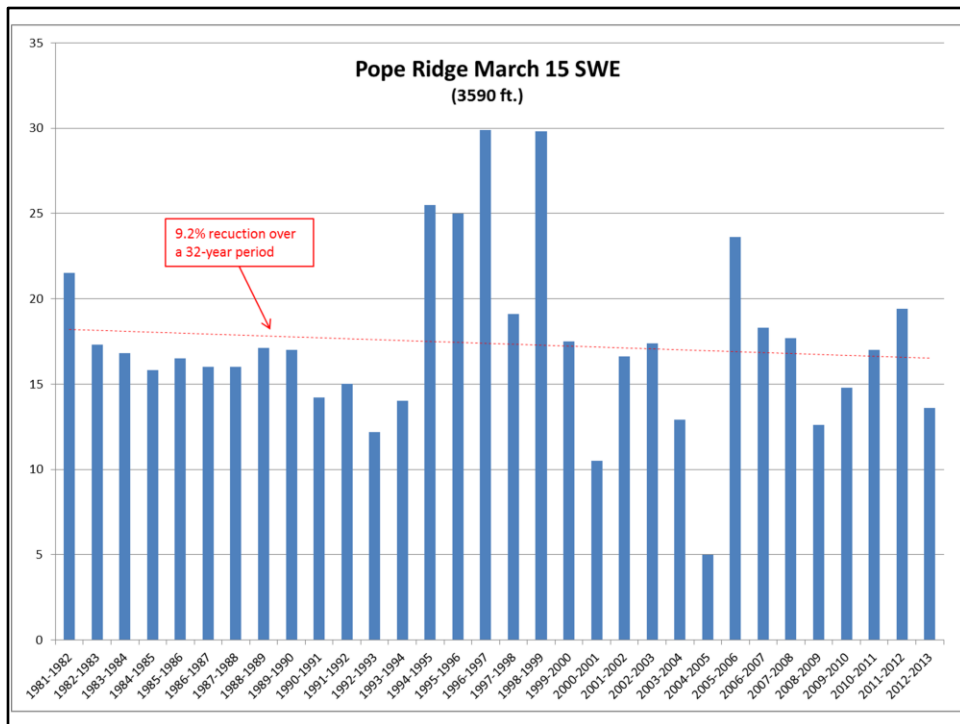
Another high-elevation site showing little change in May 1 SWE. The dates for each of these graphs are chosen by the mean peak accumulation date.



Below 1800m (5900 ft) we begin to see a stronger decline in peak date SWE. This seems to indicate earlier melting at lower elevation sites, as others in the literature have shown.



Another site below 1800m (5900 ft) showing a stronger decline in peak date SWE. This would indicate earlier melting at lower elevation sites, as others in the literature have shown.



Another site below 1800m (5900 ft) showing a stronger decline in peak date SWE. This would indicate earlier melting at lower elevation sites, as others in the literature have shown.

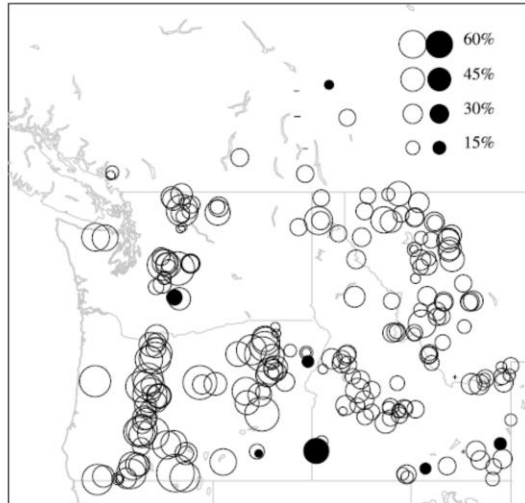
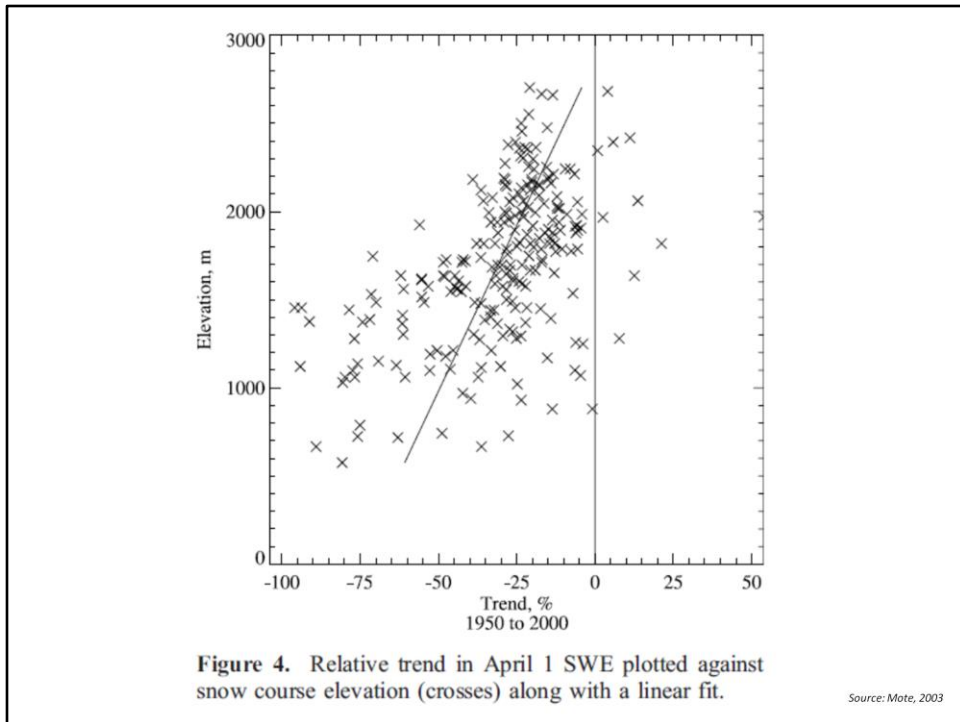


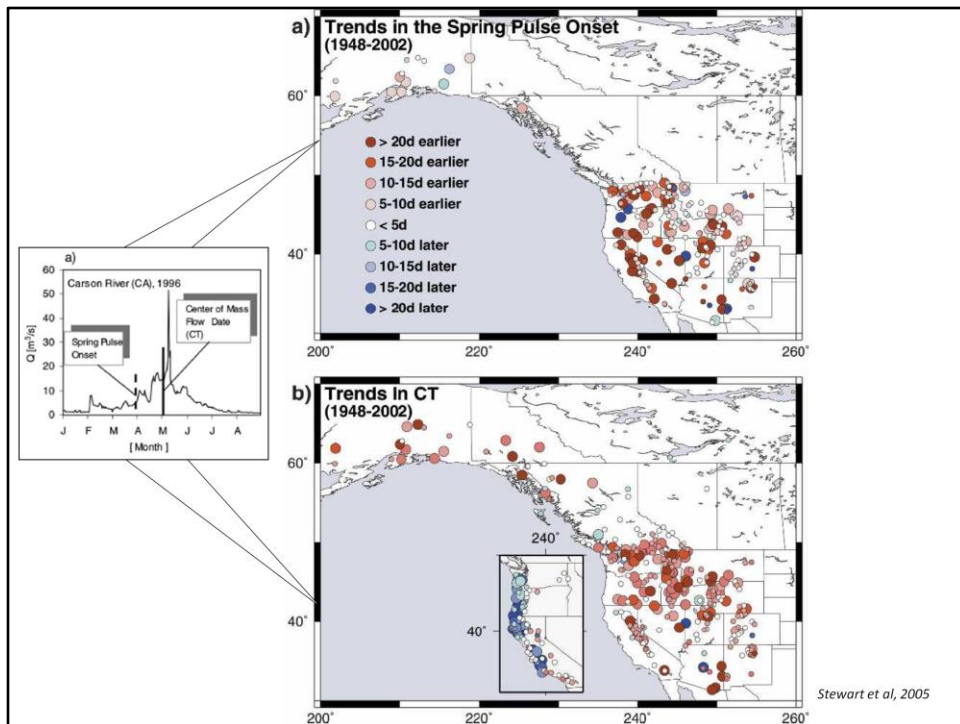
Figure 1. Linear trends, relative to starting value, in snow water equivalent (SWE) on April 1 over the period of record 1950–2000. Negative trends are shown as open circles, positive trends as solid circles; the magnitude of the trend is indicated by the area of the circle according to the legend. Trends less than 5% in absolute value are indicated by a + or – symbol.

Source: Mote, 2003

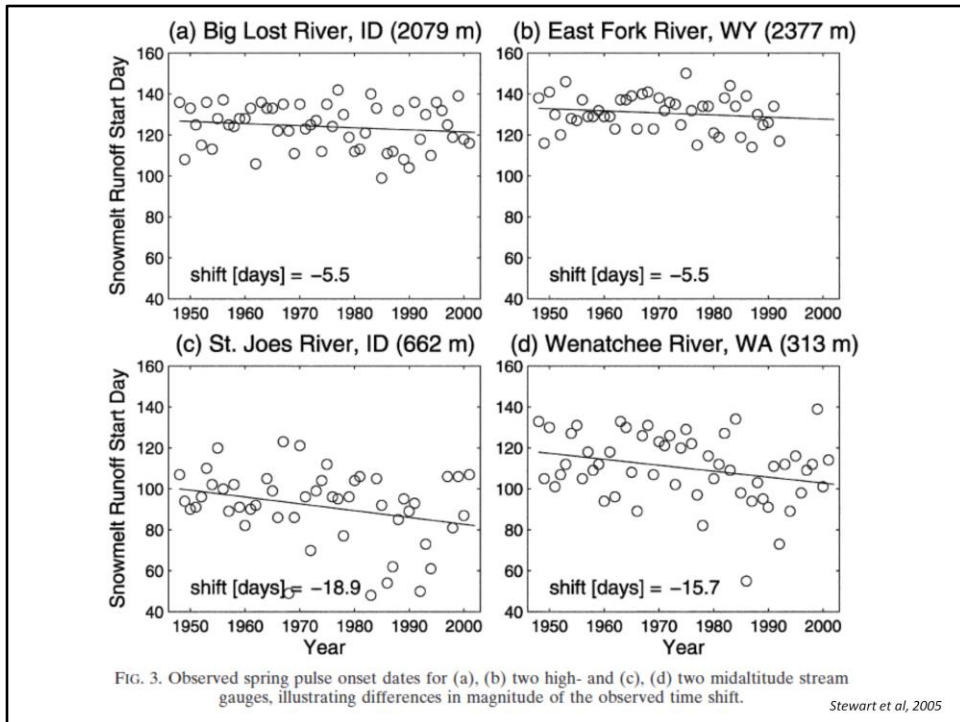
Mote (2003) evidence of largely negative trends in SWE decline on April 1st going back to 1950. Trends would likely be even stronger if data for 2001-2013 were included as these were generally earlier melting years in the PNW.



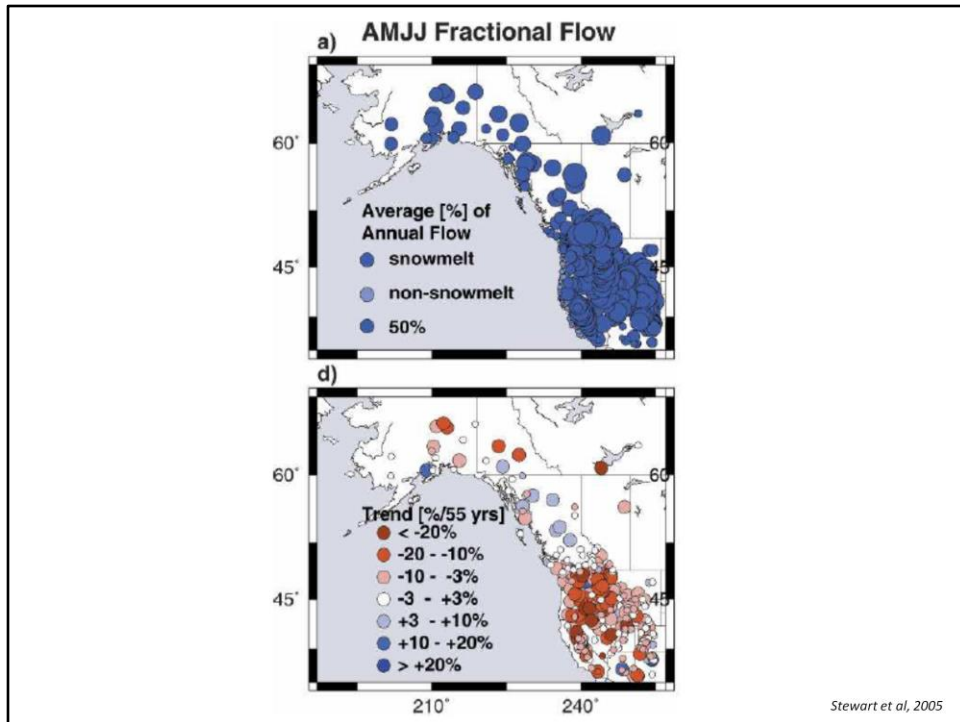
Mote (2003) evidence showing strongly negative trends, particularly below 1800m (5900 ft).



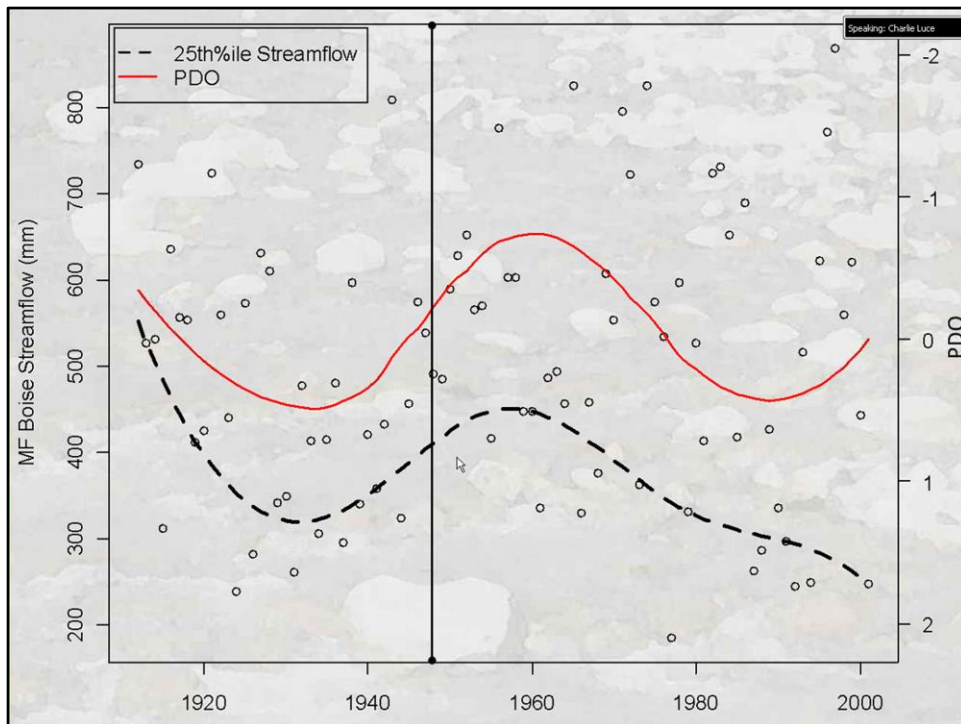
Westwide trends showing earlier onset of spring pulse (i.e., when hydrograph begins to spike upward) and center of mass (i.e., middle of hydrograph). Note that coastal rain-dominated basins are showing later trends in CT.



Shift toward earlier onset spring pulse dates (i.e., onset of runoff) is more pronounced for lower elevation sites than higher.



This shift to earlier runoff timing is very, very important in snowmelt driven watersheds across Western Nam because April-May-June-July flows comprise 50% - 80% of total annual streamflow. Earlier runoff has consequences for species and water users who rely on adequate water and habitat conditions during the low-flow periods. For both people and aquatic organisms, surviving summer low flows is a challenge.

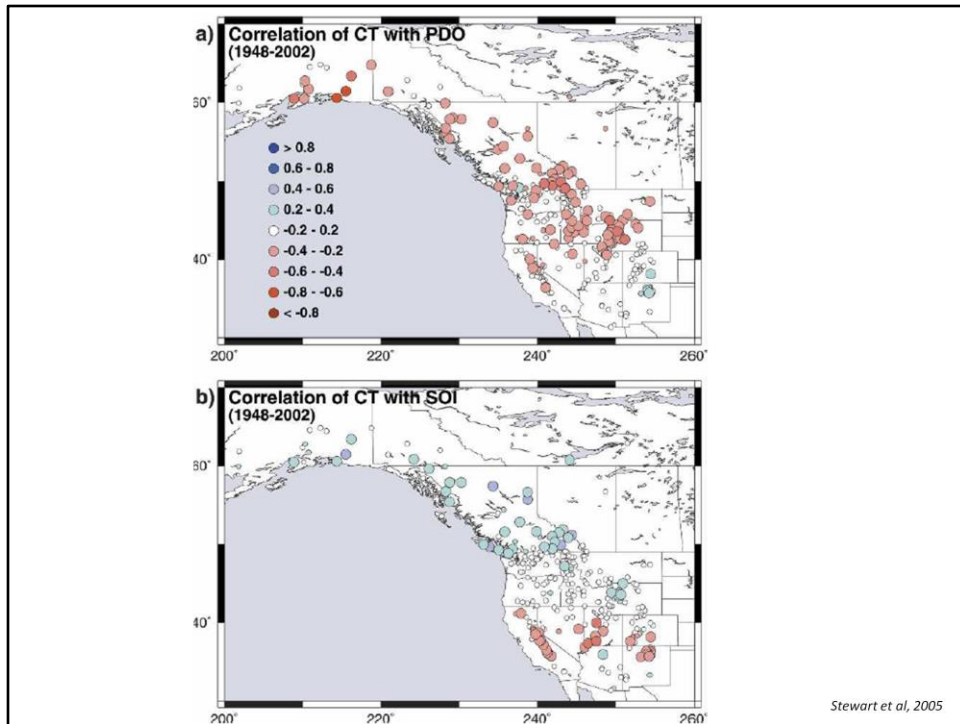


Relationship between PDO phase and (for example) the driest 25% of years.

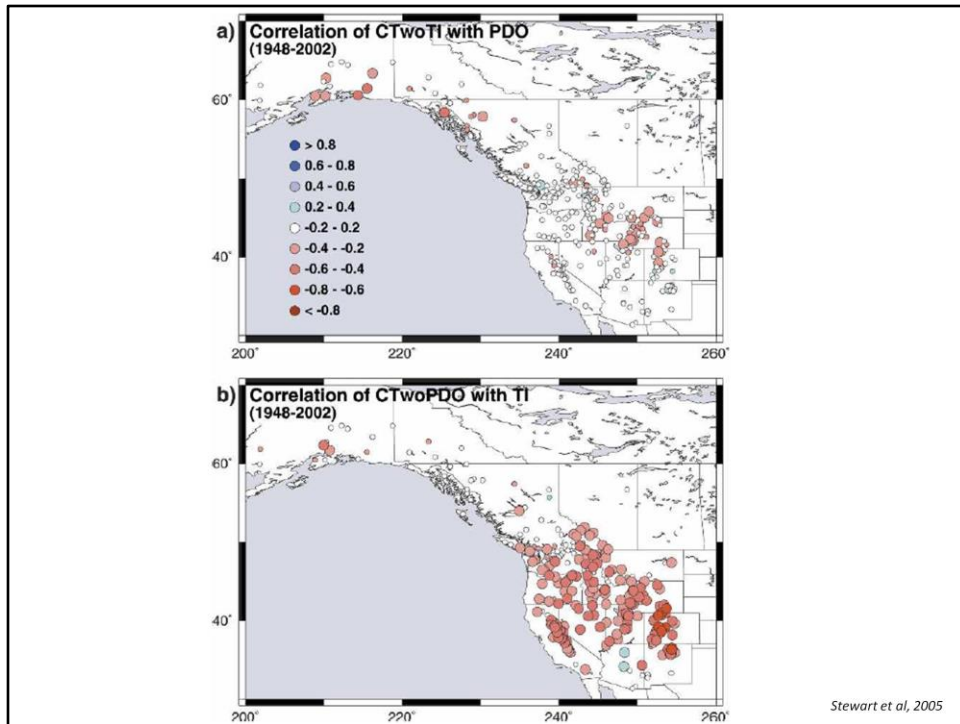
The **Pacific Decadal Oscillation** (PDO) is the leading [EOF](#) of monthly [sea surface temperature](#) anomalies (SSTA) over the North Pacific (poleward of 20° N) after the global mean SSTA has been removed, the PDO index is the standardized [principal component](#) time series.^[1]

The PDO is detected as warm or cool surface waters in the [Pacific Ocean](#), north of 20° N. During a "[warm](#)", or "positive", phase, the west Pacific becomes cool and part of the eastern ocean warms; during a "cool" or "negative" phase, the opposite pattern occurs. It shifts phases on at least inter-decadal time scale, usually about 20 to 30 years.

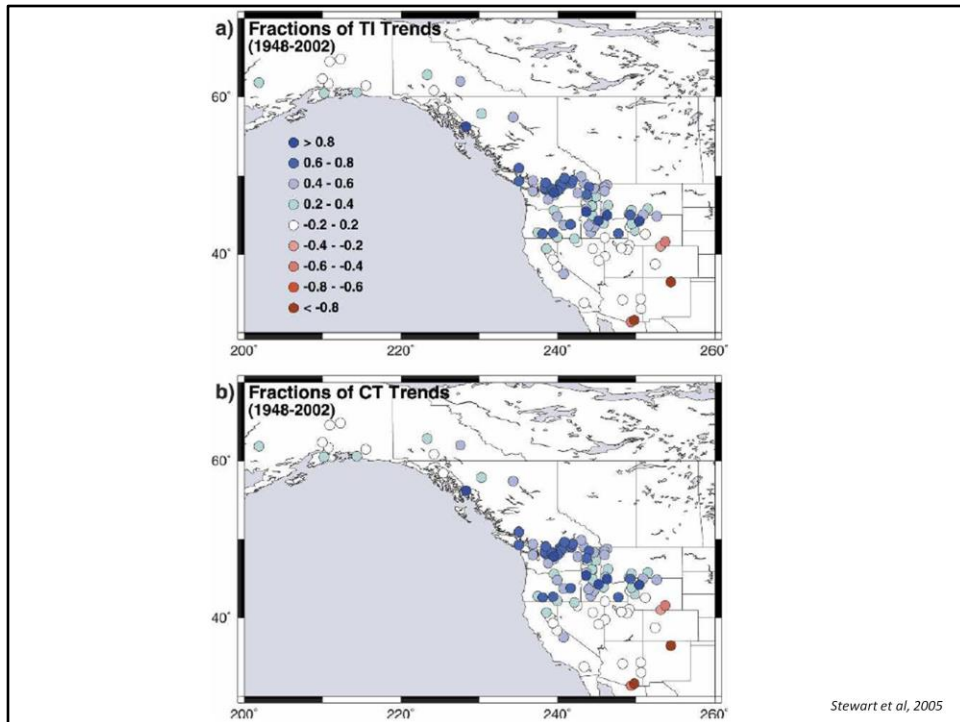
The Pacific (inter-)decadal oscillation was named by Steven R. Hare, who noticed it while studying [salmon](#) production pattern results in 1997. The prevailing hypothesis is that the PDO is caused by a "[reddening](#)" of the [El Niño–Southern Oscillation](#) (ENSO) combined with stochastic atmospheric forcing.



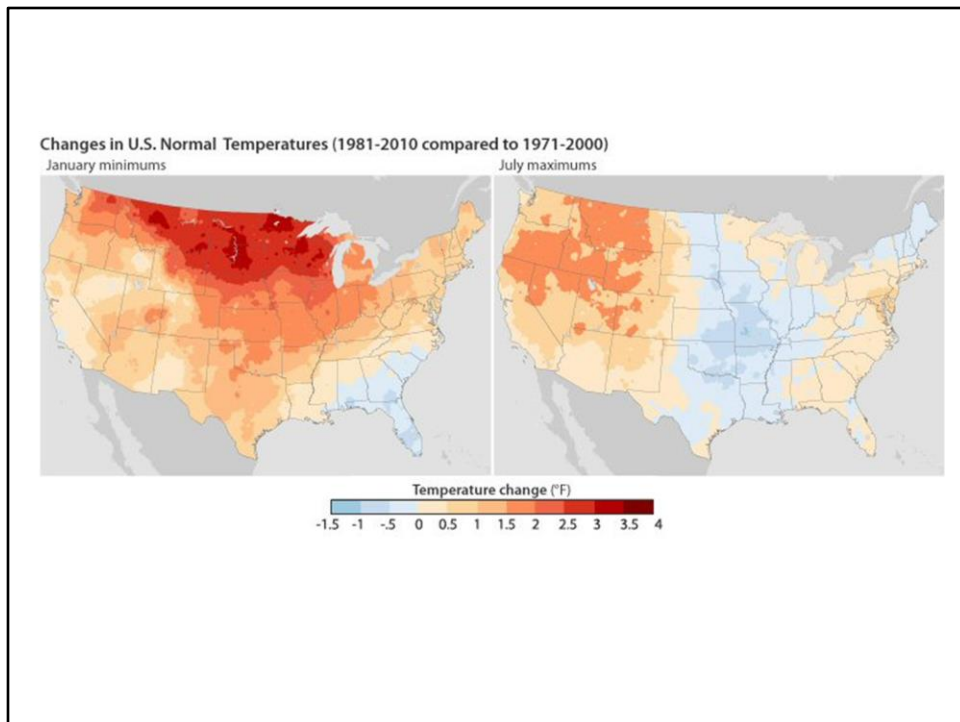
Some changes in Center Timing (CT) of runoff can be explained by PDO, particularly in the PNW where it's effect is known to be greatest (Mantua et. al. 1997) but little correlation with ENSO (Southern Oscillation Index; SOI). However, strong west-wide decreases in CT (earlier runoff) are seen irrespective of PDO phase (warm or cool).



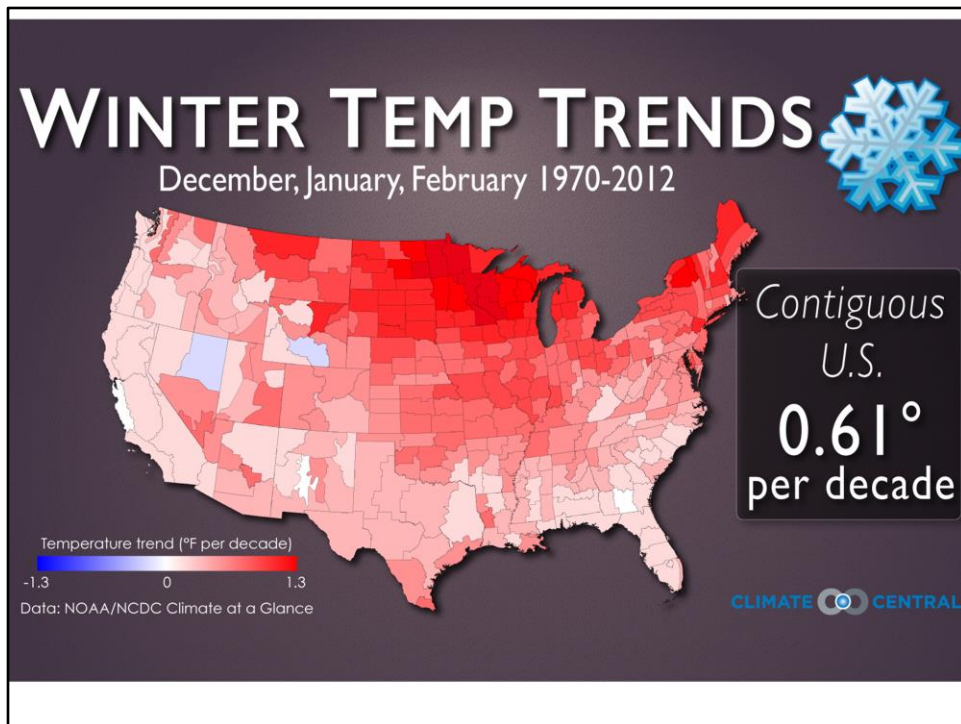
Statistical correlations of the influence of temperature (TI) and the PDO on earlier observed runoff timing. a) shows that removing the effects of warmer winter/spring temperatures but keeping PDO influences results in little correlation. b) however, when the PDO is removed, temperature alone exerts a strong control on runoff timing (CT). In other words, temperature-dependent streamflow timing variations remain when PDO influences are removed, but not vice versa.



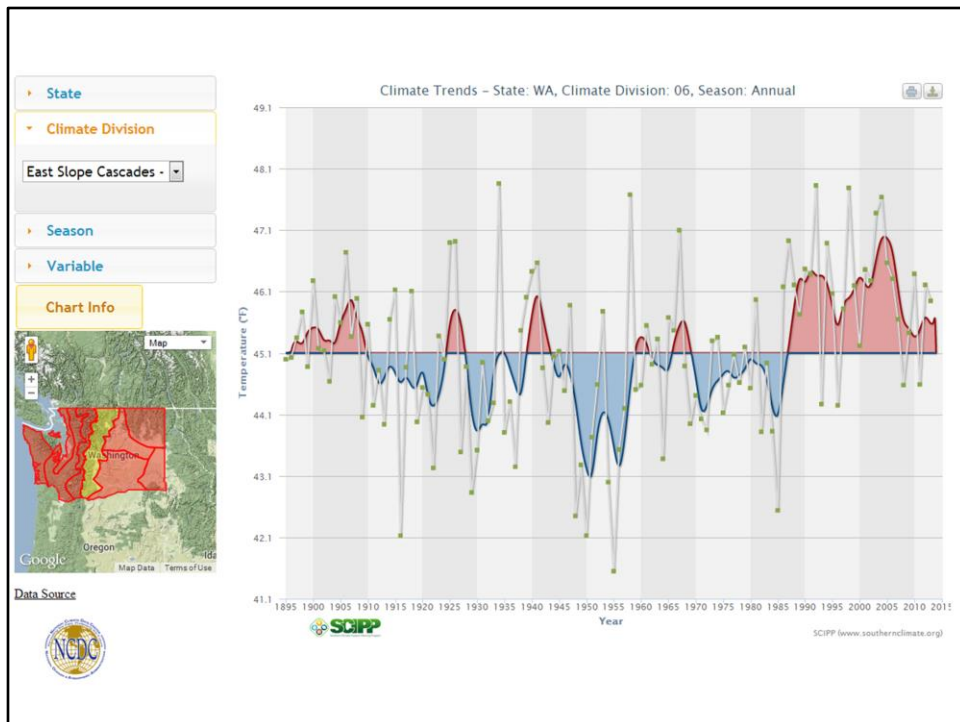
When taken together, the remainder (percentage of CT still explained by PDO) of runoff that can be explained by PDO across opposite phases (both warm and cool). In other words, PDO accounts for considerable amount of variance in runoff timing over a range of interannual to decadal, but not as much as local spring temperature (TI) variations do. The 1976-77 PDO shift accounts for less than half of the observed streamflow timing trend in most western rivers, except the PNW. Since this study, the PDO has shifted back to a cool phase and yet the streamflow timing trends are continuing unabated (evidence for broader climatological forcing).



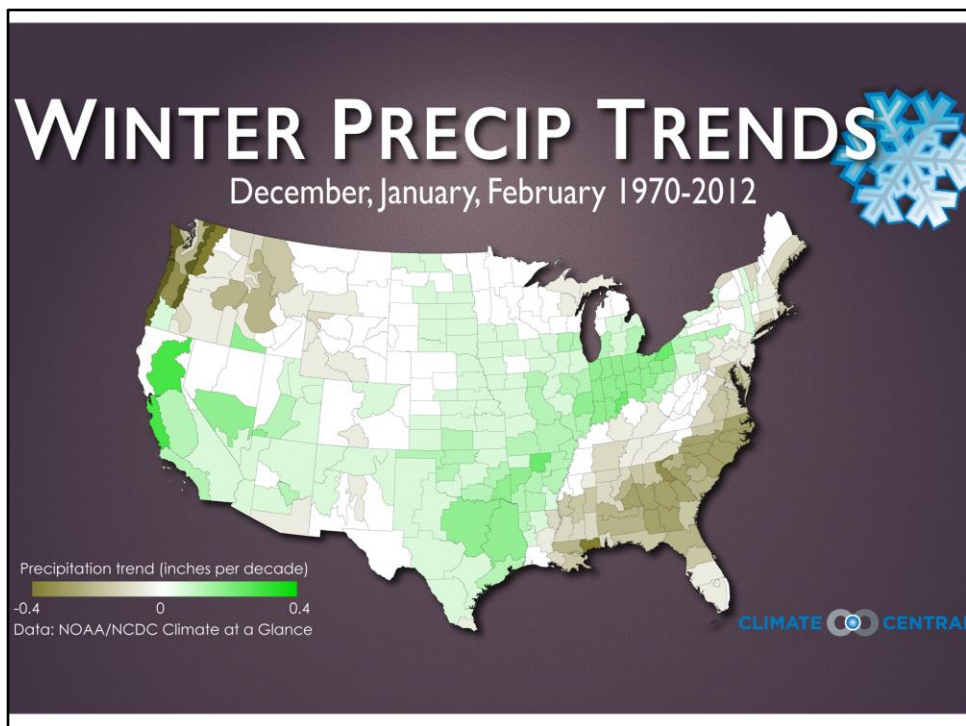
Changes in the 30-year period represent a decadal shifted baseline. A 30-year period is the standard meteorological practice employed by the World Meteorological Organization (WMO) to calculate “normals” or “averages”. It is important to note that what we consider “average” may be relative to a 30-year baseline which is numerically different from the previous. For example, in our area the 2000’s were very warm, thus shifting the new normal period to include a warmer decade than the 1970 – 2000 period. Thus, warm temperatures experienced now are closer to “average” than they would’ve been under the cooler 1970-2000 baseline. Beware the word “average” because it includes a shifting—not continuous—baseline.



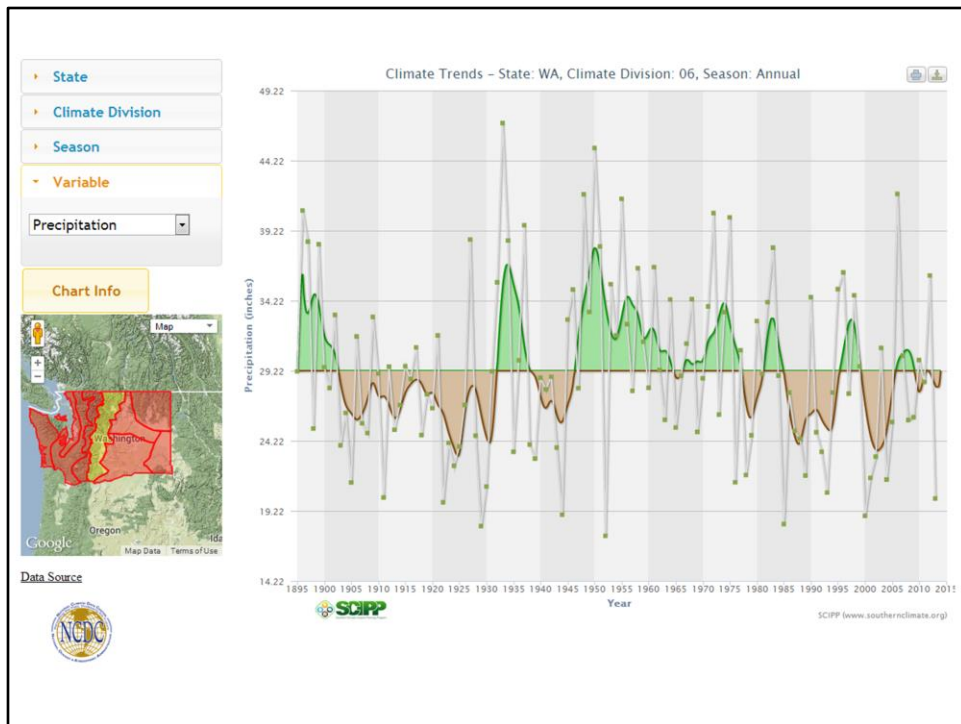
Winter temperature trends over the most recent 40-year period. Note stronger warming in the northern US and in the WA Cascades.



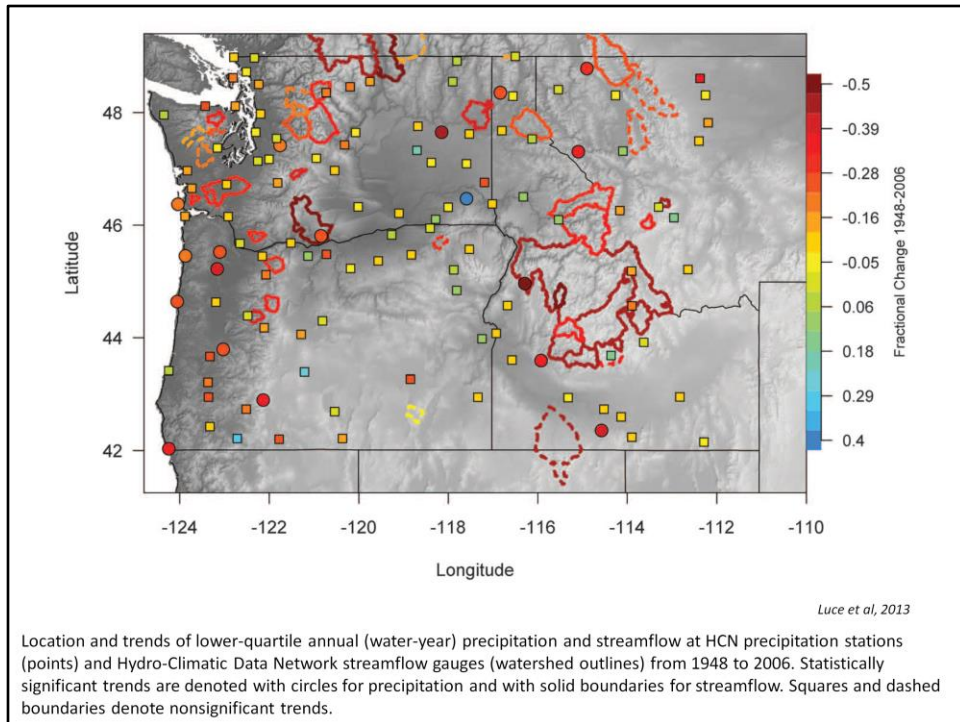
Annual changes in temperature over a longer time scale (120 years) for the eastern Cascades region. Note that considerable warming has been seen since the mid-1980s. Tree ring temperature reconstructions show that spring temperatures over the last 50 years are higher now than they've been at any time in the last 900 years. (Jones et al 2001).



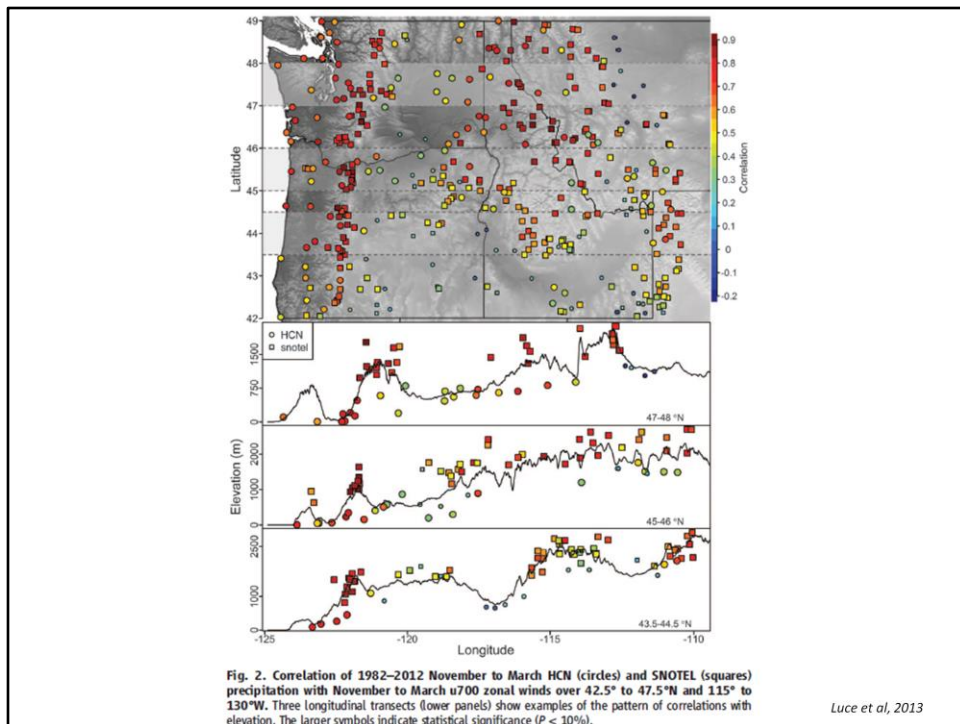
Winter precipitation trends over the most recent 40-year period. Note the sharply drier conditions in the Pacific Northwest.



Annual changes in precipitation over a longer time scale (120 years). Generally no clear pattern emerges for the eastern Cascades region.

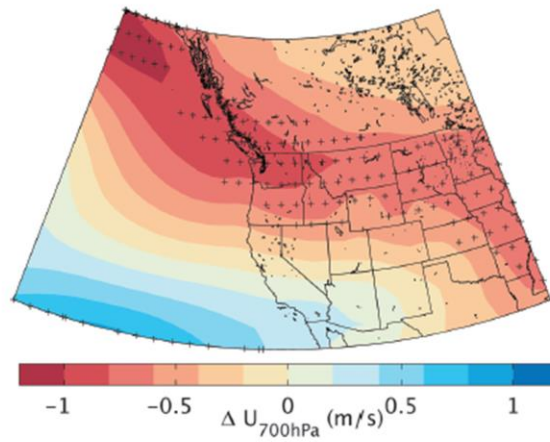


Previous studies have generally shown nonsignificant trends in precipitation over the past 50 years, despite an significant decrease in annual streamflow. This paradox has been hypothesized to be caused by increased evapotranspiration (ET) but since the precip stations do not fall within the watersheds where streamflow is measured, this makes quantitative comparison impossible. Physical calculations of energy balance (i.e., increases in ET cannot exceed incoming solar energy) have been computed and show that while there has been a small increase in ET, the observed reductions in streamflow cannot be explained by this factor alone.



Plots showing significant correlations with precipitation and 700mb (~10,000 ft) winds over the recent 30-year period on the windward side of the Cascades. Note the much lower correlations on the leeward side of mountain ranges where measured precipitation is not driven by orographic forcing. Strong evidence that orographic forcing has been reduced on the windward side of the Cascades and Rockies which better explains the reduction in annual streamflow over this timeframe (previous slide).

Fig. 4. Projected changes in November to March u700 zonal wind averaged across 24 CMIP5 models, 2071–2100 versus 1971–2000 RCP8.5. Plus signs denote areas where >80% of the models agree on the sign of the change. 20 of the 24 models predict a decrease in the PNW.



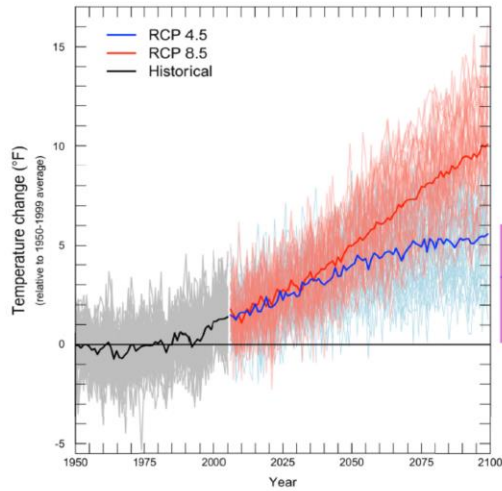
Luce et al, 2013

Late 21st century projections for 700mb (~10,000 ft) winds during the winter snow accumulation season. Strong reductions in wind speeds would have dramatic repercussions for mountain watersheds that rely on snowmelt. These physical relationships are likely not well-captured in most current climate models.

Box ES-1. Projected changes in key Pacific Northwest climate variables.

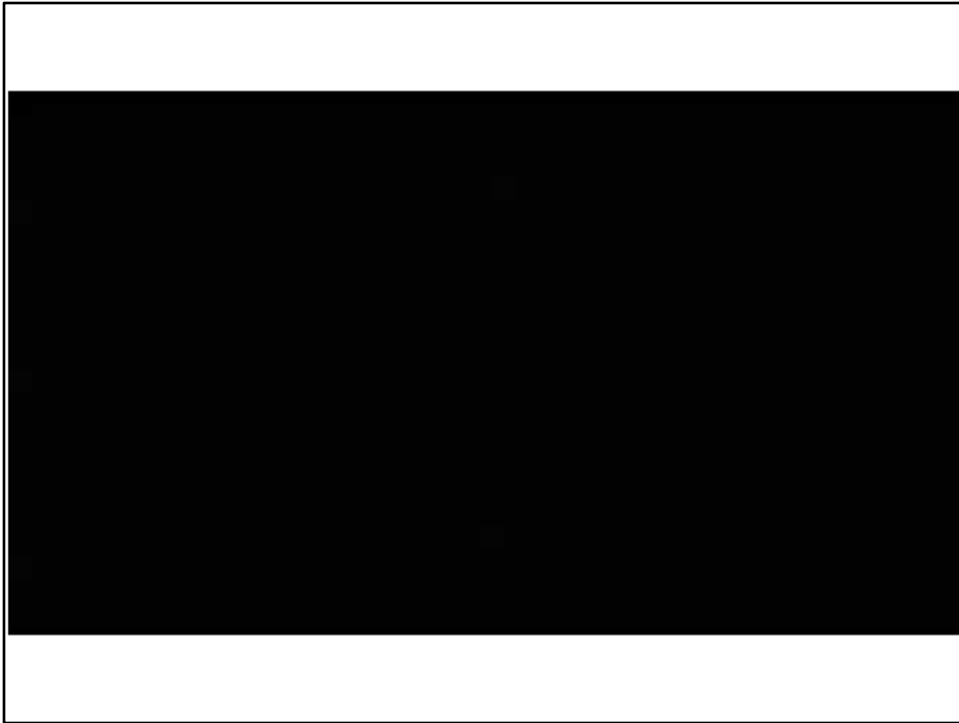
- *Average annual temperature, for 2050s:* +4.3°F (range: +2.0 to +6.7°F) for a low greenhouse gas scenario or +5.8°F (range: +3.1 to +8.5°F) for a high greenhouse gas scenario (both relative to 1950-1999).
- *Extreme precipitation, for 2050s:* number of days with more than one inch of rain increases +13% ($\pm 7\%$) for a high greenhouse gas scenario (relative to 1971-2000).
- *Average April 1 snowpack in Washington State, for 2040s:* -38 to -46% for a low and a medium greenhouse gas scenario (relative to 1916-2006).
- *Sea level in Washington State, for 2100:* +4 to +56 inches for low to high greenhouse gas scenarios (relative to 2000). Local amounts of sea level rise will vary.
- *Ocean acidity, for 2100:* +38 to +41% for a low greenhouse gas scenario and +100 to +109% for a high greenhouse gas scenario (relative to 1986-2005).

See Sections 3 and 6 for more detailed projections and additional time periods.



Univ of WA Climate
Impacts Group, 2013

Summary of most recent predicted changes to the Pacific Northwest.

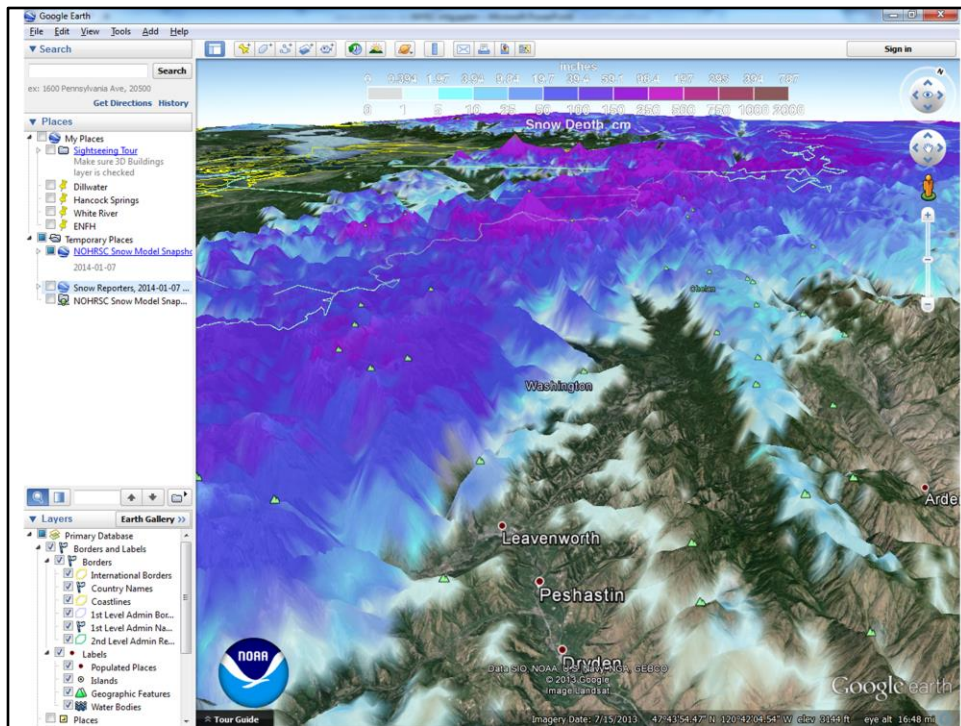


http://www.youtube.com/watch?v=_nzwJg4Ebzo

Video describing possible relationships between Arctic Amplification (warmer arctic largely due to a loss of sea ice and reduced albedo) and reduced jet stream velocities. Also results in increased jet meandering and patterns getting “stuck” longer. There is still considerable disagreement about this hypothesis in the published meteorological literature.

Conclusions

- Current drought unprecedented at lower elevations, among the worst on record at higher elevations
- Significant and ongoing reductions in SWE across all time scales—but beware the baseline for consideration of “normal”
- Continued decline in SWE on typical peak accumulation date, particularly <1800m
- Earlier runoff producing reduced predictability for humans and ecosystems
- What we’re observing can be partially explained by PDO variability but some cannot
- No significant relationship to other intra/interannual processes like ENSO
- Reasons for earlier runoff appears to be warmer spring temperatures
- Reasons for SWE decline:
 - Reductions in mountain precip (driven by reductions in orographic forcing)
 - Increases in temperatures, particularly at low and mid-elevations (<1800m)
- Potential for wide-reaching effects to hemispheric circulation due to changes in equatorial-polar temperature gradient



The National Operational Hydrologic Remote Sensing Center (NOHRSC) produces many great products now available on the web that are derived from remotely sensed (i.e., satellites) sources that cover an entire area (as opposed to SNOTEL sites which are only single points).

References

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Useful Links

Univ. of Washington Climate Impacts Group:

<http://cses.washington.edu/cig/>

UW Hydrologic Climate Change Scenarios for Columbia R Basin:

<http://warm.atmos.washington.edu/2860/>

USGS County-Level Climate Change Projections:

http://www.usgs.gov/climate_landuse/clarid/nex-dcp30.asp

Daily NRCS SNOTEL % of Average SWE and Precipitation Tables (updated daily):

<http://www.wcc.nrcs.usda.gov/ftpref/data/snow/update/wa.txt>

NRCS WA SNOTEL Statewide Map and Data:

<http://www.wcc.nrcs.usda.gov/snotel/Washington/washington.html>

Monthly WA State Climatology Newsletters:

<http://www.climate.washington.edu/newsletter/>

WA State Climatology Trends (Temp, Precip, SWE for 1895 – 2010 by region):

<http://www.climate.washington.edu/trendanalysis/>

NOAA Remotely Sensed Snow Data:

<http://www.noahrs.noaa.gov/>